



Evaluating and improving WAsP using tall wind measurements

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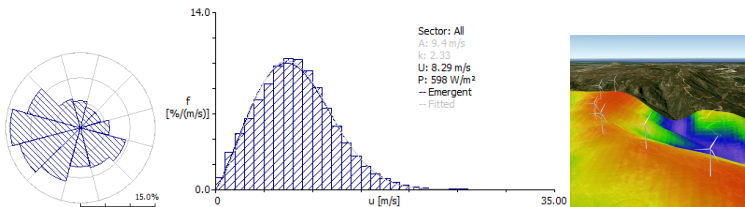
Rogier Floors Ib Troen Mark Kelly Alfredo Peña Sven-Erik
Gryning

DTU Wind Energy, Risø campus – Department of Wind Energy

EMS 2015, Sofia, Bulgaria, Energy Meteorology session

The WAsP (Wind Atlas Software Program) profile model

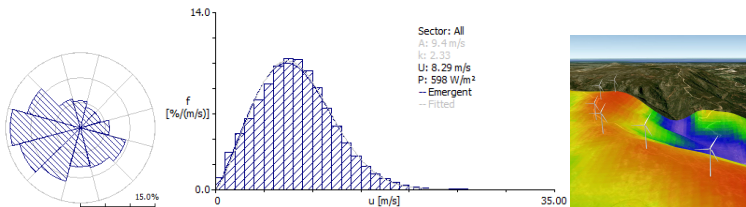
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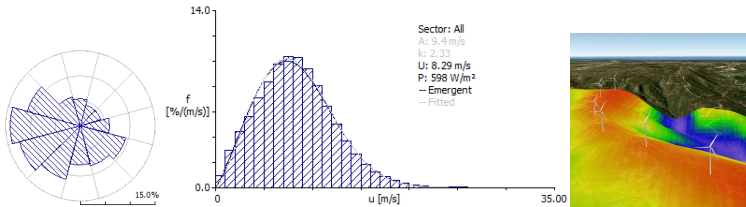
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- Contains orographic model, roughness change model, shelter model¹, wake model and vertical profile model.



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- Based on the Weibull distribution, with parameters A (scale) and k (shape).
- Contains orographic model, roughness change model, shelter model¹, wake model and vertical profile model.
- Extensively used worldwide for wind resource estimations (4500 users, 6000+ turbines sited)

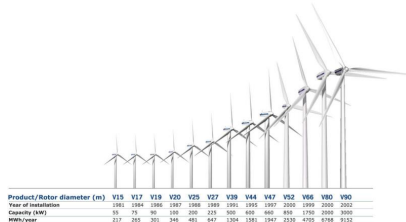


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Introduction

Motivation: 'tall' turbines reach above the surface layer and large-scale processes become more important at these heights.

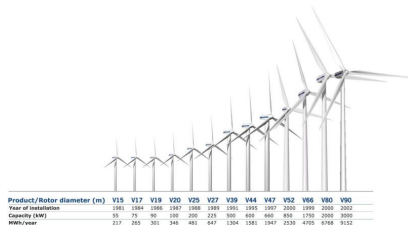
- **Improve the profile model in WAsP for both wind speed ($U = A\Gamma(1 + 1/k)$) and power density ($E = \frac{1}{2}\rho A^3\Gamma(1 + 3/k)$)**



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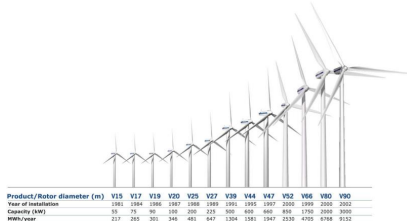
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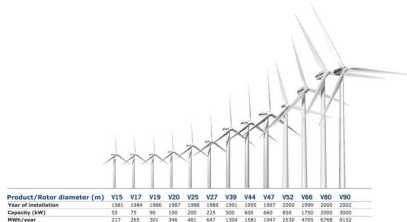
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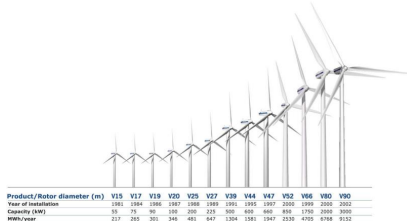
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- Implementation of baroclinicity effects from CFSR data
- Evaluation using 28 meteorological masts



Main ingredients from wind profile model in WAsP

WAsP models the effect of stability by using the logarithmic wind profile and the geostrophic drag law,

$$\frac{U}{u_*} = \frac{1}{\kappa} \left[\ln \left(\frac{z}{z_0} \right) - \psi_m \left(\frac{z}{L} \right) \right] \quad (1)$$

$$\frac{G}{u_*} = \frac{1}{\kappa} \sqrt{\left[\ln \left(\frac{u_*}{fz_0} \right) + A \right]^2 + B^2} \quad (2)$$

A and B are functions of internal stability parameter ($\mu \equiv \kappa u_* / fL$)

A first-order expansion in heatflux H is used for a u_* offset from neutral,

$$\frac{du_*}{u_*} = \frac{cg}{fT_0 c_p \rho G^2} dH \quad (3)$$

with

$$c = B \frac{dB}{d\mu} - \left(\ln Ro + \ln \frac{u_*}{G} - A \right) \frac{dA}{d\mu}, \quad (4)$$

In WAsP c is assumed constant and ≈ 2.5 . c also influences the reversal height, the height where there is maximum in the k parameter of the Weibull distribution. Consistent way of treating H_{off} and H_{rms} .

Wind lidar measurements

- G = wind lidar wind speed 950 m ($h \ll 950$ m)

Measurements described in Floors, R., Peña, A., and Gryning, S.-E. (2015a). **The effect of baroclinicity on the wind in the planetary boundary layer.**

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- u_* = iteratively solving diabatic wind profile for u_* , θ_* , q_* , L based on temperature, humidity and wind profile

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Stability $dA/d\mu$ and $dB/d\mu$ in near-neutral conditions

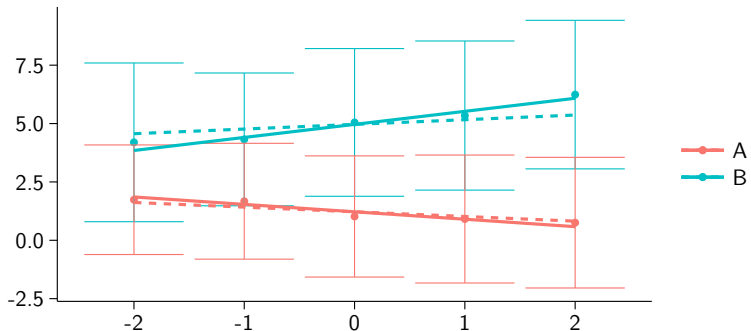


Figure: A and B as a function of stability (μ) for near neutral regime. Current linearizations used in WAsP (± 0.2) shown with dashed lines

Neutral A and B and wind direction

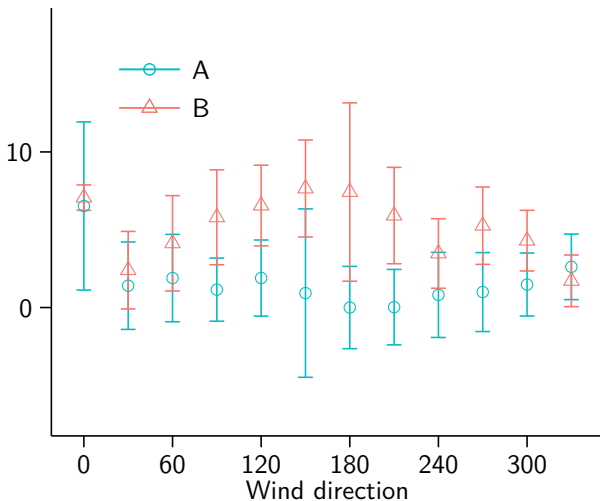
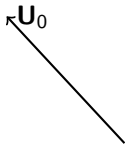
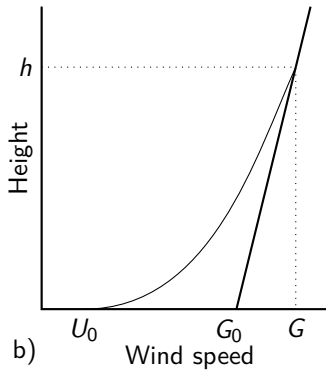


Figure: Neutral observations A and B , error bars denoting the standard deviation in each 30° bin. **Why dependent on wind direction?**

Baroclinic wind profile in theory

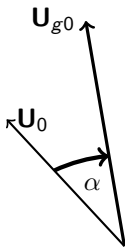


a)

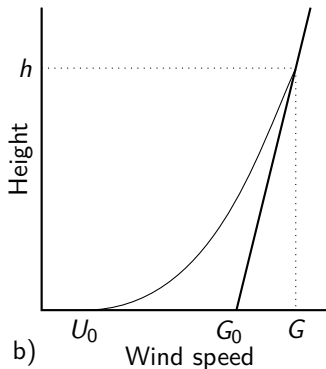


b)

Baroclinic wind profile in theory



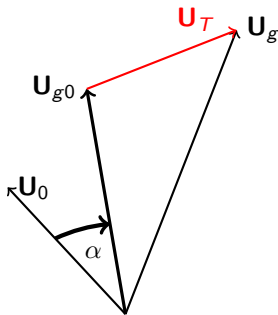
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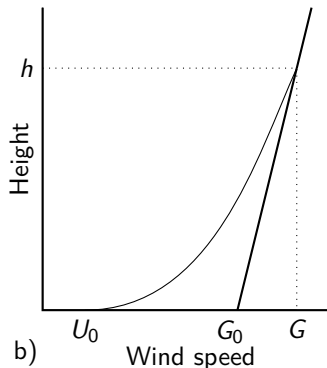
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$$\theta = \arctan(V_g/U_g) = \text{geostrophic wind direction}$$

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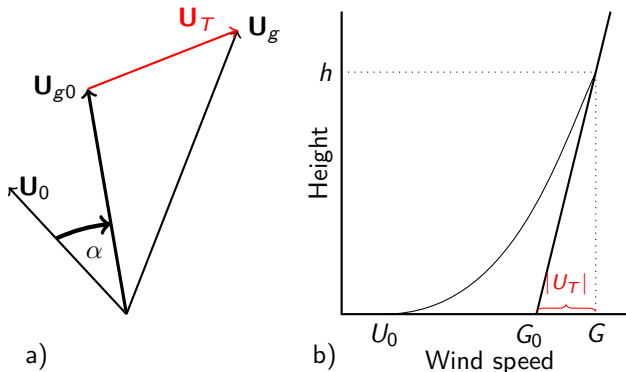


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$\theta = \arctan(V_g/U_g) = \text{geostrophic wind direction}$

$\gamma = \arctan\left(\frac{dV_g}{dz} / \frac{dU_g}{dz}\right) = \text{long-term mean thermal wind direction}$

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$\theta = \arctan(V_g/U_g) =$ geostrophic wind direction

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$M_0 = \frac{\kappa^2}{f} \frac{d|U_T|}{dz} =$ long-term dimensionless geostrophic shear

Modelling the effect of baroclinicity

A and B also function of baroclinicity,

$$A = A_0 + A' = A_0 + M_0 \cos(\theta - \gamma - \alpha), \quad (5)$$

$$B = B_0 + B' = B_0 + M_0 \sin(\theta - \gamma - \alpha), \quad (6)$$

- Eq. 5 and 6 were implemented in WAsP
- Thermal wind vector is obtained from CFSR data from geopotential height fields at 1000 and 950 hPa (± 0 –500 m)

θ = surface wind direction

γ = long-term mean thermal wind direction

M_0 = long-term dimensionless geostrophic shear

α = wind veer in boundary-layer

Baroclinicity and A and B

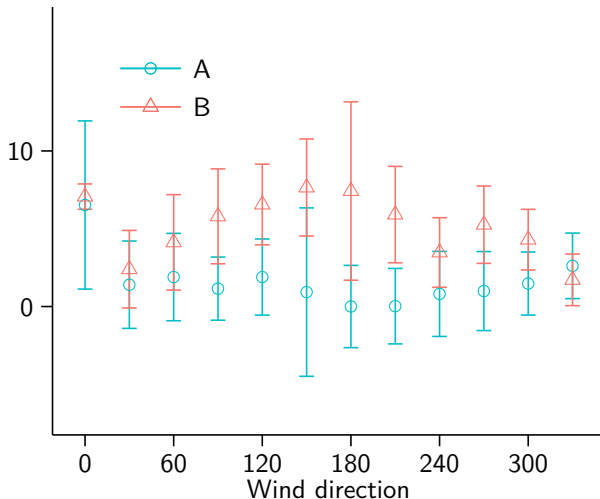


Figure: Observations A and B , error bars denoting the standard deviation in each 30° bin. The lines denote Eqs. 5 and 6 using $M_0 = 2.2$ and $\gamma = 300^\circ$, obtained from CFSR data.

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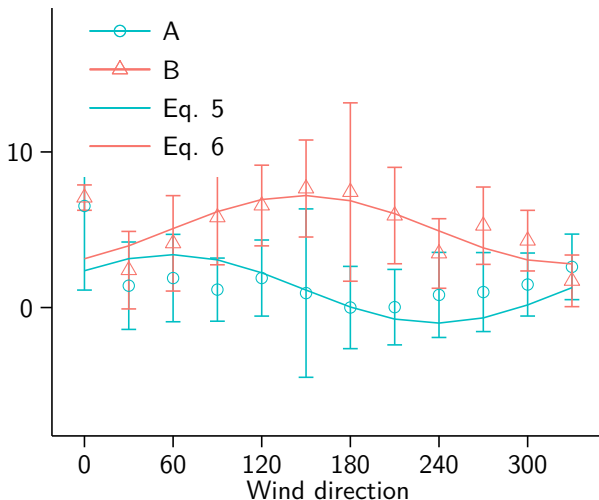


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Evaluating the models

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- ❸ Exact solutions + baroclinicity from CFSR (Troen and Petersen, 1989; Floors et al., 2015b)

Error metrics

- Mean absolute relative error, root-mean-square relative error

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- Cross predictions upwards (i.e. 10-40, 40-100 and 10-100 m)
- Exclude tall wind measurements, total of 28 sites

Sites

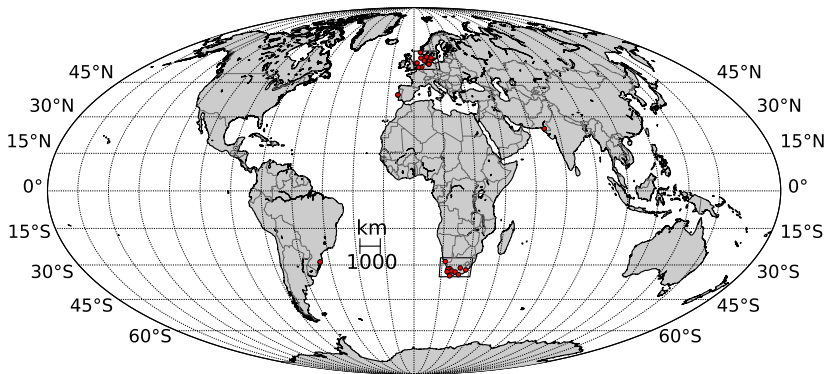


Figure: Overview of sites used for the model evaluation.

Sites

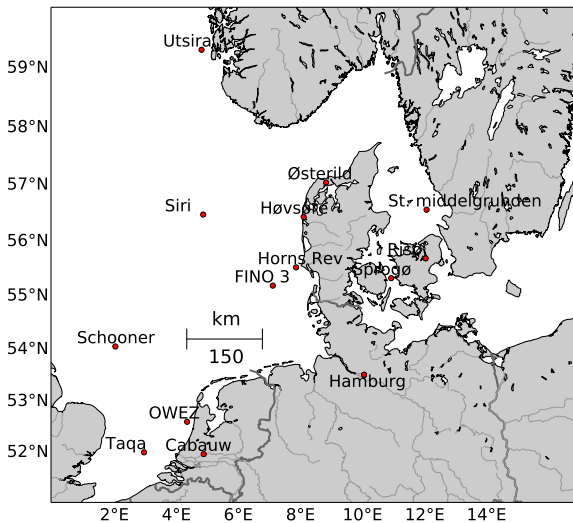


Figure: Overview of sites near Denmark.

Sites

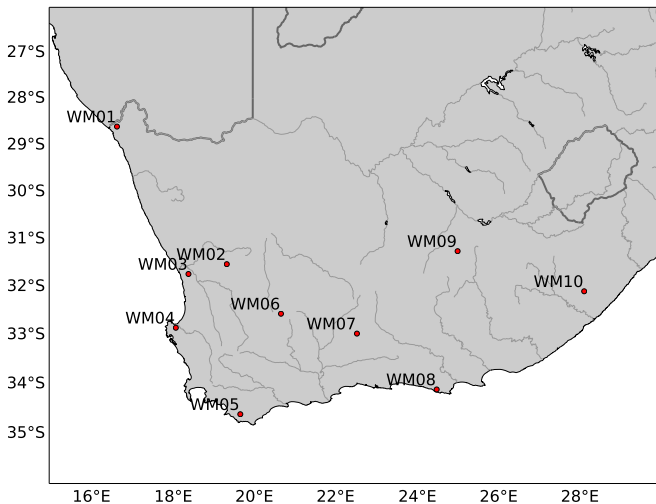
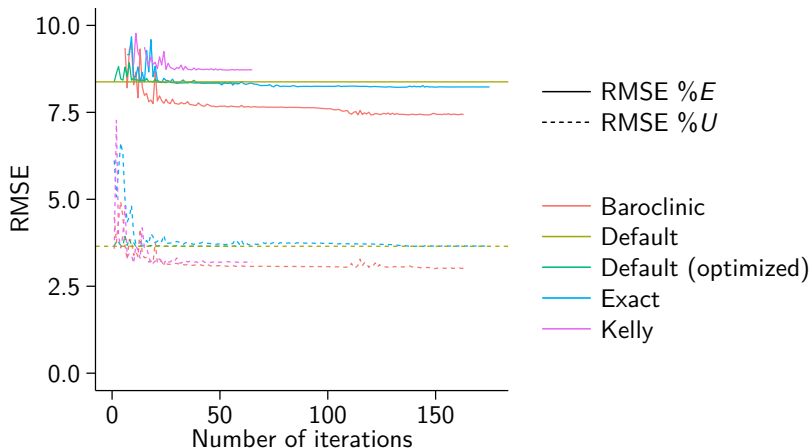


Figure: Overview of sites in South Africa.

Optimization using new default stability settings

Because the stability treatment in the model is now different, we optimized models to find values of H_{off} and H_{rms} that minimized RMSE in U and E



Conclusions

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- ② Implemented baroclinicity using just two additional parameters for each site from CFSR data
- ③ Baroclinicity + exact model approximately 10-20 % better than old model in all error metrics
- ④ Improvements most easily visible at 'tall' extrapolations

Thank you for your attention.

The study is supported by the Danish Research Agency Strategic Research Council (Sagsnr. 2104-08-0025) “Tall wind” project and internal funding. We acknowledge the BMU sponsored FINO3 mast project for the mast measurements and the Test & Measurement department for processing all measurements. We also acknowledge NCEP for the CFSR data set.

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